

ENVIRONMENTAL AND WATER RESOURCES ENGINEERING DEPARTMENT OF CIVIL ENGINEERING PORTLAND STATE UNIVERSITY PORTLAND, OREGON 97207-0751

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HYDRAULIC AND HYDROLOGIC MODEL

OF THE

FLOW CONTROL STRUCTURE

AT

SMITH & BYBEE LAKES

by

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A research project report submitted in partial fulfillment of the requirements for the degree of

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The faculty advisor approves the research project report

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1.0 **INTRODUCTION**

The Metropolitan Service District (METRO) is responsible for managing Smith and Bybee Lakes and the closure and long term management and monitoring of the St. John's Landfill. As part of the management process, METRO decided to investigate the possibility of constructing a flow augmentation structure to allow water from the Columbia River to enter the Smith and Bybee Lake system and exit the lake system via the existing flow control structure between Bybee Lake and the North Slough. The flow of water from the lakes into the North Slough would have a flushing effect which would improve the water quality of the North Slough, specifically the dissolved oxygen concentrations, to the levels of Smith and Bybee Lakes. The flow augmentation would take place during the late summer and early fall. This time period is when the North Slough experiences its worst water quality due to low flows associated with the low tidal levels of the Willamette River. This process would flush out the North Slough and possibly improve its water quality characteristics

The goal of this research was to predict the inflow from the Columbia River through a flow augmentation structure to the lake system and then to predict the outflow to North Slough.

2.0 THE SMITH AND BYBEE LAKE SYSTEM AND SURROUNDING AREA

Figure 1 shows the orientation of the Columbia River, Lower Columbia Slough, North Slough, St. John's Landfill, and Smith and Bybee Lakes. Smith and Bybee Lakes are shallow and hydraulically connected. The Columbia River is North-East of the lake system and reaches its closest point near the northern most point of Smith Lake (approximately 1600 feet). The North Slough is hydraulically connected by a flow control structure to the Lower Columbia Slough at the east end of North Slough and borders Bybee Lake along its southwestern shore. The St. John's Landfill is bordered by the North Slough, the Lower Columbia Slough, and Smith and Bybee Lakes. The Columbia River, Lower Columbia Slough, and North Slough are all tidally influenced.

3.0 **MODEL CONCEPT**

Table 1 shows water quality data for Smith and Bybee Lakes and North Slough. The most significant indicator of poor water quality is the dissolved oxygen $(D.O.)$ concentration. The lowest recorded D.O. concentration in North Slough between April 1993 and August 1994 was 2.1 mg/l on April 15, 1993. On the same day the D.O. concentrations at Smith and Bybee lakes were 10.62 and 9.98, respectively. Hence, flushing the North Slough with water from the lakes would improve the North Slough's water quality. Wells (1995) discusses in detail the impact of this flushing on North Slough water quality. Figure 2 shows the conceptual model. Figure 3 shows a flowchart of the model concept. Flow into the lake system would occur when the water surface elevation (WSE) in the Columbia River exceeds that of the lake system. The two options for the proposed inflow structure are an open channel or a closed channel (e.g., culvert). Flow into the North Slough would occur when the lake system WSE exceeds that of the North Slough which is influenced by the head in the Columbia River

ABSTRACT

computer model was created to estimate the amount of water that would enter the Smith and Bybee Lake system through a proposed flow augmentation structure connected to the Columbia River and leave the lake system through the existing flow control structure at the end of the North Slough. The model was calibrated with lake level data from the drawdown of the lakes in the Fall of 1993. Calibration included evaporation and precipitation

Lower Columbia Slough System

Figure 1: Lower Columbia Slough System

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Figure 3: Conceptual Model Flowchart

Figure 5: Open and Closed Channel Variables

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The open channel configuration between the Columbia River and Smith Lake evaluated was a trapezoidal channel. Trapezoidal shaped channels are the hydraulically most efficient shape. Figure 4 shows this configuration

Figure 5 shows the variables used in the hydraulic model for both the open and the closed channel configurations

Figure 6 shows the existing flow control structure at Bybee Lake and the North Slough. It has a two inlets for water from the lakes; one higher than the other. A flap gate on the North Slough end prevents water from North Slough from entering the lake system

Existing Flow Control Structure

Figure 6: Existing Flow Control Structure

4.0 MODEL METHODOLOGY

4.1 Objective

The objective of the model was to predict the inflow to the lake system from an open or closed channel, incorporating gravity flow from the Columbia River. The outflow from the lake system to the North Slough would then be estimated based upon the head of the lake system and the design of the existing flow control structure. An estimate of the water losses of the lake system due to evaporation and of the water gains through precipitation were also made. The model was calibrated using using meteorological and lake system water surface level data recorded during the drawdown that took place between September - November 1993.

4.2 Assumptions

The model did not account for groundwater flow into or out of the lake system. (The results presented in the water balance model below seem reasonable without additional losses or gains from groundwater.) The evaporation and precipitation estimates were based on meteorological data collected by the National Weather Service at Portland International Airport and was assumed to be representative of the Smith and Bybee Lakes area Runoff from precipitation was not included in the model

4.3 Input Data

Water surface elevation data have been collected by Portland State University (PSU) at the east end of the North Slough and in the lake system. The U.S. Army Corps of Engineers (COE) provided historical water surface elevation data for the Columbia River. Meteorological data were from the National Weather Service at the Portland International Airport. These data (file NOAAMOD.DAT), the average daily temperature, average dew point, precipitation, and average wind speed, were input data to the model, HYDRO.FOR, which calculated the precipitation and evaporation water losses and gains

4.3.1 Evaporation Model

The evaporation in a lake system can be approximated by the following equation (Gupta, 1989):

$$
E_{\text{dev}} = 0.0138e_{a}(1 - RH)(1 + 0.0098W)
$$

where

 E_{day} = evaporation per day (inches) e_a = saturation vapor pressure at the mean air temperature (mm Hg) $RH =$ Relative Humidity $W =$ wind speed (miles/day)

The saturation vapor pressure was calculated from the following equation (Linsley et al, 1982):

$$
e_a = 25.4*(-0.132579 + 0.014123*T_a - 0.000233125*T_a^2 + 2.98306*10^{-6}*T_a^3)
$$

where

 e_s = saturation vapor pressure at the mean air temperature (inches Hg)

The relative humidity, RH, can be approximated by the following equation (Linsley et al, 1982):

$$
RH = 100\left(\frac{112 - 0.1T_a + T_d}{112 + 0.9T_a}\right)^s
$$

where

 T_a = air temperature (°C) T_d = dewpoint temperature ($^{\circ}C$)

The output file from the program HYDRO.FOR, HYDRO.DAT, provided the evaporation and precipitation in inches per day. This file was used in the main model, MODEL FOR, in the SUBROUTINE EVAPPREC (Appendix D).

 $4.3.2$ Lake System

Prior studies (Fishman 1987) have established relationships between the lakes water surface elevation and volume. From these data a mathematical relationship was developed using the graphing software, SURFER, such that for any lake water surface elevation the volume and surface area can be estimated (Figure 7).

A regression analysis was performed on the area and volume of the lakes as a function of water elevation. The regression formula for volume as a function of water level was:

 $vol = -1757.08 + 10.7763 * h + 50.5127 * h²$

where

 $vol =$ lake volume (acre-feet) h = water surface elevation (feet MSL)

The regression equation for lake surface area as a function of water elevation was:

 $area = -4151940 - 9228770 * h + 3632870 * h² - 292929 * h³ + 7310.35 * h⁴$

where

area = lake surface area $(f²)$ h = water surface elevation (feet MSL)

These regression equations were determined using GRAPHER's curve fitting functions.

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433 Proposed Flow Augmentation Structur

The location for the flow augmentation structure was initially chosen at the location of the shortest distance between the Columbia River and the lake system. The distance was approximately 1600 feet. The model had the options of using an open channel or a closed channel (culvert). Common variables for each configuration were the invert elevation of the channel at the Columbia River, invert elevation where the channel meets the lake system, length, slope, Manning's roughness coefficient (a function of the channel material), and an estimate of a discharge coefficient to apply to the one way flow gate where the channel terminates at the lake system.

The open channel had the following variables: channel side slope, bottom width. The open channel was assumed to be constructed of concrete. The closed channel had the following additional variables: type of material (i.e., concrete or corrugated metal) and diameter.

The open channel flow calculations were made in the SUBROUTINE OPENC in the program MODEL FOR. The fundamental equation applied was Manning's equation for steady, open channel flow (Gupta 1989);

$$
Q=\frac{1.486}{n}AR^{2/3}S^{1/2}
$$

where

Q= discharge (cubic feet per second [cfs])

 $A = \csc$ sectional area of discharge (f_1^2)

 $R =$ Hydraulic Radius (ft)

 $S =$ channel slope (ft/ft)

The closed channel flow calculations were based on fundamental culvert design equations. In this model the three types of flow used were types 3, 4, and 5. Figures 8-10 show the possible flow situations. For Type 3 culvert flow, the flow was estimated by the following equation (Gupta 1989):

$$
Q = C_d A_{\text{exil}} \sqrt{2g(H+z-h_{\text{index}}-h_{\text{exil}}-h_{\text{cubic}})}
$$

where

Q= discharge (cubic feet per second [cfs])

 C_d = discharge coefficient

 A_{exit} = cross sectional area of discharge (ft²) at exit

 $g =$ gravitational acceleration (ft²/s)

h = height of Columbia WSE above the channel exit invert

 $h_{\text{index}} = \text{index head loss}$

 h_{crit} = exit head loss

 h_{cutoff} = head loss through the culvert

Type 5 Flow (H/D > 1.2)

Figure 10: Type 5 Culvert Flow

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For Type 4 culvert flow, the flow was estimated by the following equation (Gupta 1989):

$$
Q = C_d A_c \sqrt{\frac{2g(h - h_{ext})}{1 + 29C_d^2 n^2 L / R_c^{4/3}}}
$$

where

 Q = discharge (cubic feet per second $[cfs]$)

 C_d = discharge coefficient

 A_c = cross sectional area of culvert (f_t^2)

 $g =$ gravitational acceleration ($\frac{f}{f}$ /s)

 h = height of Columbia WSE above the channel exit invert

 h_{unit} = exit head loss

 $n =$ Manning's roughness coefficient for the culvert material

 $L =$ length of culvert (f)

 R_c = hydraulic radius of the culvert

For Type 5 culvert flow, the flow was estimated by the following equation (Gupta 1989):

$$
Q = C_d A_c \sqrt{\frac{2g(h-D)}{1 + 29C_d^2 n^2 L / R_c^{4/3}}}
$$

where

Q = discharge (cubic feet per second [cfs])

 C_d = discharge coefficient

 A_c = cross sectional area of culvert (f_1^2)

 $g =$ gravitational acceleration ($\text{ft}^2\text{/s}$)

 h = height of Columbia WSE above the channel exit invert

 $D =$ diameter of the culvert (ft)

 $n =$ Manning's roughness coefficient for the culvert material

 $L =$ length of culvert (ft)

 R_n = hydraulic radius of the culvert

The determination of which type of flow to use was based on the water surface elevations at both ends of the culvert.

4.3.4 Flow Control Structure Between North Slough and Bybee Lake

The existing flow control structure, Figure 6, consists of a 62.5 ft. long corrugated metal pipe (CMP) with a 60 inch diameter. At the North Slough end is a flap gate that prevents water from the North Slough from entering the lake system. There is a high flow overflow segment, an adjustable weir, and a canal gate. The adjustable weir has a minimum elevation of 8.4 ft. mean sea level (MSL) and water reaches the weir through a 36 inch diameter grated intake with an invert elevation $(i.e.)$ of 6.9 ft. MSL.

The water reaches the canal gate through a 30 inch diameter g/rated intake with an invert elevation of 5.5 ft. MSL. For lake levels below 5.5 ft. MSL there will be no flow through the structure.

4.3.4.1 Adjustable Weir

The adjustable weir was modeled as a weir with two end contractions (Figure 11). The following equation was used to estimate the flow over the weir (Gupta, 1989):

$$
Q = \frac{2}{3} C_d \sqrt{2g} (L - 0.2H) H^{3/2}
$$

where.

 $Q =$ discharge (cfs) C_d = discharge coefficient $g =$ gravitational acceleration ($\frac{\hat{\pi}^2}{s}$) $H =$ upstream head above the weir (ft)

In the development of this equation it was assumed that the upstream velocity was zero. Head losses due to the intake structure will be accounted for in the weir discharge coefficient. The adjustable weir has a maximum elevation of 13.4 ft. mean sea level (MSL) and a minimum elevation of 8.4 ft. MSL.

4.3.4.2 Canal Gate

The canal gate (Figure 12) consists of a 30 inch circular opening that is covered by a circular plate. The circular plate can be raised to open the canal gate The canal gate was modeled as sluice gate The following equation was used to estimate the discharge through the canal gate (Gupta, 1989):

$$
Q = C_{\rm g} A \sqrt{2gh}
$$

where.

 $Q =$ discharge (cfs) C_{g} = discharge coefficient $g =$ gravitational acceleration (ft²/s) $h =$ upstream head over the gate (ft)

As with the adjustable weir, the head losses due to the intake structure will be accounted for in the canal gate discharge coefficient

5.0 ESTIMATE OF FLOWS FROM COLUMBIA RIVER TO SMITH LAKE

5.1 Vaiiables Used

 $\frac{\partial \theta_{\rm in}}{\partial \theta_{\rm in}}$ $\hat{g}_{\rm gas}^{1/2}$ $\frac{\partial \mathcal{L}_\text{max}}{\partial \mathcal{L}_\text{max}}$

For both open and closed channel configurations of the proposed flow augmentation structure slope of 1 foot per 1000 feet was chosen. This is a very mild slope and initially chosen to minimize the vertical drop over the length of the channel (1.5 ft drop over a length of 1500 ft.). For the open channel the bottom width was varied from 10 - 20 ft. and a constant side slope of 1:1 was used. For the closed channel a diameter of 6 - 12 feet was used. The Manning's roughness coefficients were chosen as mid points in their ranges for the types of materials. The invert elevation of both channels at the Columbia River was initially set to 9 ft MSL. This was set because 1600 ft from the Columbia River would be at the 9 ft MSL contour in Smith Lake (see Figure 13). The initial discharge coefficients used for both the weir and the canal gate were 0.6 which was based upon typical coefficients for those types of structures (Gupta 1989).

The water surface elevation data for the Columbia River and North Slough was from December 3, 1992 - February 11, 1993.

5.2 Results

It was immediately apparent that the closed channel configuration was impractical as evident in the very small rise in the lake levels No significant flow was reaching the lake system This was due to the much smaller cross-sectional area of the discharge for a circular channel as opposed to a trapezoidal channel at the same invert elevations The model was run for the open channel configuration with varying bottom widths and entrance invert elevations And the canal gate was varied for closed, half open, and fully opened configurations. The weir was set at the lowest point, 8.4 ft. MSL. The model predicted very little change in flow through the canal gate from half open to fully open It also predicted very little net inflow from the Columbia River during this time period. Appendix A contains some representative graphs of the results (note that the time axis is in Julian day and corresponds to the time period of from December 3, 1992 - February 11, 1993).

Historical Columbia River peak daily head data were obtained from the U.S Army Corps of Engineers (COE). Figure 14 show the data mean daily peak head covering the years 1973-1990. These data show that the Columbia River water surface elevation is very low during the thne period of late summer and early fall. The historical data indicate the impracticability of using a flow augmentation structure that relies on gravity flow during low-water periods

6.0 CALIBRATION OF HYDRAULIC MODEL BETWEEN BYBEE LAKE AND NORTH SLOUGH

6.1 Calibration Data

METRO performed a drawdown test on the lake system from September - November 1993. Besides taking water quality data, water levels in the lakes and North Slough were recorded. The North Slough water surface elevation data from the same time period were also used. The results of the drawdown test and final model calibration are shown in Figure 15. Figure 16 shows the final model output of predicted lake level and discharge through the existing flow control structure

Figure 13: Smith and Bybee Lake Bathymetry

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Figure 15

Page 18

Figure 16

Figure ¹⁷ shows the water losses through the flow control structure and through evaporation and the water gain through precipitation

62 Calibration

The model was calibrated by adjusting the discharge coefficients for the weir and canal gate When the lake system water surface elevation was below the weir elevation, the outflow was only from the canal gate. During calibration it became evident that the weir's elevation was closer to 8.2 ft MSL rather than the 8.4 ft MSL specified in the engineering plans. Metro had surveyed the structure in 1993. During the survey it was noted that the corner of the structure where the weir is located was approximately 0.2 feet below the specified plan elevation. The model was adjusted to reflect a weir elevation of 8.2 ft MSL. The discharge coefficients used in the final calibration were 0.65 and 0.07 for the weir and canal gate, respectively. The discharge coefficient for the weir was in the typical range (Gupta 1989). The discharge coefficient for the canal gate seemed low, and probably reflect the entrance losses gate losses and exit losses Other losses occurred due to plant matter and debris build up on both of the entrance grates

70 CONCLUSION

The proposed flow augmentation structure was impractical due to low Columbia River water surface elevations during the period of late summer and early fall and the shallow lake system. The low water surface elevation in the Columbia River did not provide enough head for gravity flow into the lake system and then out through North Slough Any flow augmentation during this time period would require pumping water

The calibration results do not indicate that any significant amount of water is entering or exiting the lake system through groundwater interactions during the drawdown period.

A program, PUMP.FOR (provided in Appendix C), calculates the net volume, length of time, and increase in water surface elevation of the lake system for a known pumping rate. If water could be pumped into the lake system it could allow for the flushing of the North Slough.

The total amount of precipitation during the drawdown period was 1.6 inches. The estimated loss through evaporation was approximately 7.5 inches. Although it was difficult to predict the amount of rainfall, the amount of rainfall relative to the evaporation was comparatively small. If the model is run for a representative year of meteorological data, it could be used to reasonably predict the drawdown of the lakes during low precipitation periods The outflow through the structure could also be used as input data to a water quality model of the North Slough.

Figure 17

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References

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APPENDIX A

Initial Model Results

APPENDIX

Program Code for Main Model

```
icol - Col Rv time index
        hcol - Col rv head
ŵ
        hby - Bybee Lake head
                                                idg - diffg time ndex
\mathcal{R}hns - North slough Head
                                                ins - NS time index
\starPROGRAM model
         IMPLICIT DOUBLE PRECISION (A-H, O-z)
        Double precision jday, jns
        common jns (9000), hns (9000)
        real hday(122), time(5000), head(5000)
        REAL*8 mano, lo, kc, lc, manc, gatelevel
         character channeltype, pipetype, gatestatus
\mathfrak{m}OPEN (UNIT = 21, FILE = 'ens06.dat', STATUS = 'OLD')
        OPEN (UNIT = 24, FILE = ' colhd2.use', STATUS = ' OLD')
\mathbb{C}^*STATUS = 'OLD')OPEN(UNIT = 25, FILE = 'program.in',
C
        open (unit = 26, file = 'program.dat', status = 'unknown')
       open(unit = 27, file = 'average.dat', status = 'unknown')
C
         open (unit = 28, file = 'hydro.dat', status='old')
\Phi* polynomial coefficients vol = -1757.08 + 10.7763*hth + 50.5127*hth**2
    derived from the application of the "grapher" best fit 2nd degree
\mathcal{M}polynomial to the Fishman results.
女
\Phia=-1757.08b=10.7763c = 50.5127\Phiprint *, 'Cd weir:'
         write(*,*) ' '
         read *, cdweir
         write(*, *) ''
         print *, 'Cd gate:'
         write(*, *) '
                        \mathcal{L}read *, cdgate
         write(*, *)\mathbf Cread (25,100) elevc, slopec, cdc, lc, diac, kc, elevo, slopeo, cdo,
\mathbf Clo, bo, dyo, dxo, mano, channeltype, pipetype, gatestatus,
\mathbf{C}+gatelevel
\mathbf{C}+format(6f10.4,/,8f10.4,/,a1,9x,a1,9x,a1,9x,f10.2)
100
\frac{d}{d\lambda}WRITE (*, *) ''
\mathbf{C}WRITE (*, *) ' What type of channel would you like to use? '
\mathbf{C}WRITE (*,*) ' Closed or open channel (C-closed, O-open)?'
\mathbf{C}\mathbf{C}WRITE (*, *) ''
        read(*,10) channeltype
\mathbf{C}C10format(a4)\frac{1}{2}if (channeltype .eq. 'c') then
\mathbf{C}WRITE (*, *) ' Do you want a concrete pipe or a corrugated '
\mathbf{C}write (*, *) '
                            metal pipe c-concrete, m-metal?'
\mathbf{C}WRITE (*, *) ''
\mathbf{C}read (*,10) pipetype
C.
        endif
\mathbf{C}if (pipetype .eq. 'c') manc = 0.013\mathbf{C}if (pipetype .eq. 'm') manc = 0.024\mathbf C\starWRITE (*, *) ''
\mathbf{C}\mathbf{C}WRITE (*, 9)
        FORMAT(' Is the canal gate opened? (Y - yes, N - no).')
C<sub>9</sub>WRITE (*, *) ''
\mathbf{C}
```

```
\mathbf CREAD (*, 10) gatestatus
C
s set gatestatus = yes
\mathbf{C}gatestatus = 'y'\tilde{\mathbf{x}}TE (*,*) ' '<br>(gatestatus .EQ. 'y') THEN<br>WRITE (*,*) ' How much is it raised? ( Fully open is 2.5 ft.<br>WRITE (*,*) ' (Welf of the area is 1 75 ft.)'
           WRITE (*, *) ''
\mathbb{C}^*IF (gatestatus .EQ. 'y') THEN
               (gatestatus .EQ. 'y') THEN<br>WRITE (*,*) ' How much is it raised? ( Fully op<br>WRITE (*,*) ' (Half of the area is 1.75 ft.<br>WRITE (*,*) ' '
C
\mathbb{C}\mathbb{C}WRITE
READ 40 GATELEVEL
C
c
               WRITE (*, *) ''
\mathbb{C}^*\mathbf CENDIF
\starqatelevel = 2.5\pmb{k}\mathbf Cif (gatestatus .eq. 'n') gatelevel = 0.0\mathbf{r}\mathbb{C}^{\mathbb{C}}WRITE (*, 15)citCher = 0.15 = 0.15<br>c15 = FORMAT ('Enter the heigth of the adjustable weir in feet.', /,<br>c = 0.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1.15 = 1
                 Use a decimal point.', /,
                  The minimum weir height elevation is 8.4 ft and the
maximum is 13.4 ft
\mathbb{C}+ \rightarrow\mathbb{C}c + \prime maximum is 13.4 ft.\prime)<br>
39 READ (*, *) \prime \prime<br>
40 FORMAT(F5.2)
          WRITE (*, *) ''
         FORMAT(F5.2)weir = 8.2\mathbf{r}IF (WEIR .LT. 8.4 .OR. WEIR .GT. 13.4) THEN
\mathbf{C}\mathbf{C}WRITE (*, 41)FORMAT (\prime, \prime The number you entered was outside the operational', \primeC41+ , ' parameters. Try again.', \sqrt{ }\overline{\mathbf{u}}GO TO 39
C
a.
          ENDIF
\bullet* read in the Col. Rv., North slough head data, their respective times in
\pmb{\hat{\pi}}Julian Day, and the diffq file with rspective time.
\mathbf{r}hcol - Col rv head jcol - Col Rv time index<br>hby - Bybee Lake head jdg - diffq time index
ŵ
                                                           \ldots jdq - diffq time index
\Phihns - North slough Head jns - NS time index
Ŵ.
蠢
* read in the Bybee Lake head from the same file, 'model.dat', as the diffq
*<br>c301 format(9x,f8.3,13x,f5.2)
c302 format(f11.3,20x,f10.3,10x,f10.3)c303 format (f8.3, 5x, f5.2)do 304 i=1, 9000read(21,*, end=305) jns(i), hns(i), crap<br>jns(i)=jns(i)-365
304 enddo
305 insend =i-1read in data from hydrologic model, hydro.for. Julian day, evap.,
\star\mathbf{\hat{x}}and prec
\starread ind hydro.dat julian days and subtract 365 to math other data files
\star
```

```
\frac{d}{d} 306 i=1, 122
          read28 end333 hdayi crap1crap2
          hday(i) = hday(i) -365306 enddo
333 ihydro =i-1rewind(28)\Phi\PhiC.
         do 314 i=1, 7000
           read(23,302, end=306) jdq(i), hby(i), diffq(i)\mathbb{C}^*c314 enddo
c306 idqend =i-1\Delta\mathbb{C}do 324 i=1, 7000
           read(24, 303, end=307) jcol(i), hcol(i)\mathbb{C}^*c324 enddo
c307 icolend =i-1*<br>308 write(*,*) ' '<br>write(*,*) 'The data files have the following start/end times:'
          write(*, *) '
          write(*, 309) jns(1), jns(insend)write(*, 312) hday(1), hday(ihydro)write(*, 310) jdq(1), jdq(idqend)C
         write(*,311)jcol(1), jcol(icolend)<br>write(*,*)C
\Phi309 format ('North Slough:', 2f11.3)<br>c310 format ('diffq/hby: ', 2f11.3)
c310 format('diffq/hby:c311 format (' Col. River: ', 2f11.3)<br>312 format (' Hydrologic Data:', 2f11.
        format ('Hydrologic Data:', 2fll.3)
会
\frac{1}{24}Set the start time for the program.
\pmb{\ast}if(jns(1) .ge. jcol(1) .and. jns(1) .ge. jdq(1))start = jns(1)
\Gamma_{\rm tot}^{\rm eq}if(jns(1) .ge. jcol(1) .and. jns(1) .ge. jdq(1))start = jns(1)<br>if(jcol(1) .ge. jns(1) .and. jcol(1) .ge. jdq(1))start = jcol(1)
\mathbf{C}if(jcol(1) .ge. jns(1) .and. jcol(1) .ge. jdq(1))start = jcol<br>if(jdq(1) .ge. jcol(1) .and. jdq(1) .ge. jns(1))start = jdq(1)
         if(jdq(1).ge. jcol(1).and. jdq(1).ge. jns(1)) start = jdq(1)<br>write(*, 313) start
C.
c write (*,313) start<br>313 format //, start time: ', f10.3)
         if(jns(1) .ge. hday(1)) start = jns(1)
C
\mathbf Cif (hday(1) .ge. jns (1)) start = hday(1)C
         write(*, 313) start\Phistart = 986.625黄
\starSet the end time
\Phiif(jns(insend).le.jcol(icolend).and.jns(insend).le.jdq(idqend))
c
C
              end = jns(insend)if (jcol (icolend).le.jns (insend).and.jcol (icolend).le.jdq(idqend))
\mathbf C\mathbf{C}end = \text{jcl}(icolend)if(jdq(idqend).le.jcol(icolend).and.jdq(idqend).le.jns(insend))
\mathbf C\mathbf Cend = idq(idqend)+ end = jdq(10)<br>write(*,313)end
\mathbf C*<br>314 format ('end time: ',f10.3, /)
          if (jns (insend).le.hday(ihydro) end = jns (insend)
```

```
if (hday(ihydro). le. jns (insend)) end = hday(ihydro)write(*, 314) endÙ.
         end = 1052.490\mathbf C\starUse a 60 minute time increment (in Julian day)
\tilde{\mathbf{v}}\pmb{\ast}1 = hcol\mathbf{\hat{x}}variable to interpolate:
   n:2 = hns\pmb{\ast}\pmb{\mathrm{t}}3 = diffq\mathbf{\hat{x}}4 = hby\pmb{\ast}returned variable
   z:dt = 1./24.the first lake level is 9.13 ft msl
         hth = 9.13write(*, 375) hthformat \binom{7}{1} h1 = ', f6.2)
375
         total = (end - start)*24iend = int (total)\tan = \text{start}{sumout = 0.0}sumin = 0.0do 400 i=1, iend
         flowin=0.0outflow=0.0call inter (time, 1, hc, icolend, insend, idqend)
J,
           call inter (tim, hn, insend)
            call inter(time, 3, dif, icolend, insend, idqend)
C.
素
\mathbf{k}calculate outflow
\starIF (GATESTATUS .EQ. y) THEN
\mathbf{C}CALL CANALGATE (hth, QGATE, gatelevel, cdgate)
\mathcal{L}ELSE
          QGATE = 0.0\mathbf C\mathbb{C}^2ENDIF
\pmb{k}★
         IF (HN .GE. hth .OR. hth .LT. 5.5) outflow = 0.0IF(HN .LT. hth) CALL WEIRFLOW (hth, QWEIR, WEIR, cdweir)
\bulletIF(HNSLOSS .GE. 13.4 .AND. HNSLOSS .LT. hth) then
         CALL CULVERT (HNSLOSS, hth, QCUL)
         else
         qcu1 = 0.0ENDIF
         outflow = QGATE + qweir + QCULcalculate the inflow through the proposed structure
```

```
Ŕ
  c - closed, o - open\starif (channeltype .eq. 'c') then
\overline{z}call closed(flowin, hth, hc, elevc, slopec, cdc, lc, diac,
\mathbf Ckc, manc)
\mathbf C÷
\tilde{\mathbb{C}}endif
\star\mathbf Cif (channeltype .eq. 'o') then
           call openc \tilde{f}lowin, hth, hc, elevo, slopeo, lo, bo, dyo,
а.<br>Ч
÷,
                       dxo, mano)+<br>endif
\mathbb{C}^*\Phiing<br>San
        flowin = flowin + dif\Phi* convert cfs to acre-ft for time increment dt
\mathbf{d}dvin = flowin*1.983471*dt\mathbb{C}^2dv = outflow*1.983471*dt\Phicall subroutine to read in hydrologic data for evaporation and
       precipitation
\mathbb{C}\mathbb{C}call evapprec (tim, hth, evap, prec, dt, hydvol)
          dv = dv + hvdvolquadratic equation for smith/bybee lake vol vs elevation in ac ft vs ft msl
\Phiusing the coefficients a, b, c.\Delta tvol = a + b*hth + c*hth**2
\Delta\sigmanflow is considered positive
\mathcal{R}\frac{1}{2}dv = dvin - dvoutC.
          volnext = vol - dvC
豪
          time\times1m+365write(26, 401)tim, hth, outflow, evap, prec
       format (f8.3, 2f8.4, 2f7.5)
401
          \tan = \tan + \text{d}t -365
\Phihth=((-b/c)+sqrt((b/c)**2)-4.*(a-volnext)/c)/2.
\mathbf{C}if(hth.lt.8.4) stop\mathbf{r}calculate the averages of the flowin and flowout
\Phi\Deltasum = sumout + outflowsumin = flowin + suminC
\mathbf{r}400 continue
\Phiend = real (iend)avqout = sumout/endavgin = sumin/end\mathbf C\Delta r\Deltaformat(f5.0, 3x, f8.4, 3x, f4.2)505
        format(f9.3,3x,f8.4,3x,f9.8,3x,f9.8)
510
         write (27,20) channeltype, pipetype, gatestatus, gatelevel,
\mathbf{C}weir, avgout
C
         format (' channeltype: ', a4, /,' pipetype: ',a4, /,
C<sub>20</sub>
```

```
' canalgate open: ', a4, /, ' gatelevel: ', f4.1, /,
                  canalgate open: ',a4,/, ' gatelevel: ', f4.1,/,<br>weirheight: ', f5.2,/, ' avg. outflow: ', f7.2,/
\mathbf C+\mathbf{C}weirheight: ', f5.2,/,<br>avg. inflow: ', f7.2,/
\mathbb C\starwrite (27,100) elevc, slopec, cdc, lc, diac, kc, elevo, slopeo, cdo,
Ċ
                      lo, bo, dyo, dxo, mano, channeltype, pipetype, gatestatus,
\overline{\mathbb{Z}}gatelevel
đ
       \ddot{+}\starstop
END
         subroutine evapprec (tm, hd, ev, pr, dt, hydvol)
         IMPLICIT DOUBLE PRECISION (A-H, O-z)
 polynomial coefficients asurf (surface area)
         aa = -4151940bb = -9228770cc = 3632870dd = -292929ee = 7310.35rewind(28)600 real(28,*) day, evapor, precip
          day = day -365.
          if(int(tm).eq.int(day))thenevap and pr are in inches - convert to feet
\mathbf Cev=evapor*dt/12pr = precip * dt/12else
          go to 600
          endif
* hyd = hydrologic loss & gain
          hydev-pr
           hyd=ev-pr<br>asurf=aa + bb*hd + cc*hd**2 + dd*hd**:
                 =aa + bb*<br>ee<mark>*</mark>h d**4
  + ee*h d**4<br>hydvol = volume of water (ft2) lost to evap & gained through precip.
          hydvol=hyd*asurf
convert to acre-ft
          hydvol = hydvol/43560return
          end
         SUBROUTINE WEIRFLOW (HDBY, DISCH, WEIRHT, cdweir)
         IMPLICIT DOUBLE PRECISION (A-H, J-m, O-z)
         PARAMETER (PI = 3.14159)cdweir = 1.0Ċ
\PhiQ1: Flow over the adjustable weir (sharp crested.)
ŵ
   Q2: Flow over the weir structure itself (sharp crested.)
\PhiQ3: Flow over the canal gate (broad crested.)
ŵ
   Q4: Flow over the overflow structure (sharp crested.)
\mathbf{r}\starCase la
\star\bulletIF (HDBY .LE. WEIRHT) THEN
           Q1 = 0.0
```

```
02 = 0.003 = 0.0Q4 = 0.0<br>DISCH = Q1 + Q2 + Q3 + Q4END IF
        IF (HDBY .GT. WEIRHT .AND. HDBY .LT. 13.4) THEN
        Q1 = 2./3.*cdweir*4.*(2.*32.2)**0.5*(HDBY-WEIRHT)**1.5
        Q2 = 0.003 = 0.0Q4 = 0.0DISCH = Q1 + Q2 + Q3 + Q4ENDIF
\Delta* Case 1b:
        IF (HDBY .GT. WEIRHT .AND. HDBY .GT. 13.4) THEN
        01 = 2./3.*cdweir*4.*SQRT(2.*32.2)*(13.401 - WEIRHT)**1.5
        Q2 = 2./3.*cdweir*5.75*(2.*32.2)**0.5*(HDBY-13.4)**1.5<br>Q3 = 0.004 = 0.0DISCH = Q1 + Q2 + Q3 + Q4ENDIF
 Case 1c:
Ů
        IF (HDBY .GT. 13.5) THEN
        Q1 = 2./3.*cdweir*4.*(2.*32.2)**0.5*(13.401 - WEIRHT)**1.5
        Q2 = 2./3.*cdweir*5.75*(2.*32.2)**0.5*(HDBY - 13.4)**1.5
        Q3 = 0.385*(2*32.2)**0.5*5.25*(HDBY - 13.5)**1.5Q4 = 0.0DISCH = Q1 + Q2 + Q3 + Q4ENDIF
* Case 1d:
        IF (HDBY .GT. 13.8) THEN
        Q1 = 2./3.*cdweir*4.*(2.*32.2)**0.5*(13.401 - WEIRHT)**1.5
         Q2 = 2./3.*cdweir*5.75*(2.*32.2)**0.5*(HDBY - 13.4)**1.5
        Q3 = 0.385*(2.*32.2)**0.5*5.25*(HDBY - 13.5)**1.5Q4 = 2./3.*0.62*PI*6.*(2.*32.2)**0.5*(HDBY - 13.8)**1.5DISCH = Q1 + Q2 + Q3 + Q4ENDIF
\mathcal{R}RETURN
        END
SUBROUTINE CANALGATE (BYHEAD, QG, DIST, cdgate)
        IMPLICIT DOUBLE PRECISION (A-H, J-m, O-z)\mathbf{r}This subroutine will estimate the flow through the canal gate.
\Phiŵ
   The canal gate is modeled as a sluice gate. The flow is derived
\mathbf{R}by applying the Bernoulli Equation
\frac{1}{24}The distance, DIST, is measured from fully opened.
\frac{1}{N}So, 2.5 ft. is fully closed and 0.0 is fully open.
\star\mathbf{r}PARAMETER (PI = 3.14159)cdgate is the estimated discharge coefficient for the canal gate
```

```
experience indicates this is low 4/17/94
      cdqate = 1.0\mathbf CD = 2.5 - DISTA = \text{PI} * (1.25**2 - (D/2)**2)if (byhead .le. 5.5) then
        OG = 0.0else
        QG = cdgate*A*(2*32.2*(BYHEAD-4.75))**.45endif
        RETURN
        END
  SUBROUTINE CULVERT (S, BY, QC)
        IMPLICIT DOUBLE PRECISION (A-H, J-m, O-z)
   This subroutine estimates the discharge through the structure when
     the head in the North Slough is high enough to result in a Type 4culvert flow condition with minor head losses due to the 60 inch
     pipe and the flapgate
   CD = discharge coefficient.
   HYDRAD = hydraulic radius of the pipe.N = Manning's Coefficient.
   L = length of pipe.KM = minor head loss coefficient due to flap gate. (Assume = 0.45)
∴÷ke ∴
        PARAMETER (PI = 3.14159)CD = 0.7G = 32.2A = PI/4.*25.
        HYDRAD = 5./4.R = HYDRADN = 0.013L = 60.0KM = 0.45if(by .le. s) thenqc = 0.0else
        QC = CD*A*SQRT(2*G*(BY - S)/(1 + 29*CD**2*N**2*L/R**1.333 ++ CD**2*KM) )endif
\PhiRETURN
        END
subroutine closedqcol hdby hdcolelev slopecd ldiakman
\mathbb{Z}IMPLICIT DOUBLE PRECISION (A-H, J-m, O-z)
C.
       real*8 k, 1, losses, manC.
       parameter(pi = 3.14159)\mathbf{C}parameter(g = 32.2)\mathbb{Z}^d\Phi* elev = invert elevation at entrance
* elbyb = invert elevation at exit
* variables follow Gupta: Hydrology and Hydraulic Systems; pg. 656
   h12 entrance loss
   h23 = pipe loss
```

```
\starhgate exit loss
\star\mathbb{C}if(hdcol .le. hdby .or. hdcol .le. elev)then
\mathbb{C}qcol = 0.0C
             goto888
\tilde{c}endif
\mathbb{R}^n\mathbf Cif(dia .eq. 0.0) then<br>write (*,*) ' dia = 0.0'C
\overline{\phantom{a}}stop
\mathbf Cendif
\mathbb{C}bigh = hdcol - elevif (bigh. le. 0.0) then\mathbb{R}^2\mathcal{L}^{\mathcal{A}}qcol = 0.0\mathbf Cgo to 888
\mathbb{Z}endif
\frac{1}{2}radius = dia/2.
\mathbf{C}z = slope * L<br>h1 = bigh + z
\mathbf{C}iong<br>San
          elbyb = elev - zgamma = hdby + elbybC.
\mathbb{C}tfour = elbyb + dia\frac{\partial \mathbf{u}_0}{\partial \mathbf{w}^2}if(hdby .lt. elbyb)then
a.
             h4 = 0.0\mathbf Celse
\frac{1}{2}h4 = hdby - elbyb\tilde{\mathbb{C}}endif
          ratio = biqh/dia\mathbf C\mathbf Chgate = 0.45*Cd**2\mathbf{r}\hat{\mathbf{x}}ratio less than or equal to 1.2
\mathbf{R}с
          if (ratio .le. 1.2) then
蠢
\mathbf{r}Type 3 culvert flow - assume V1 = 0.0- assume flush connection at entrance (k = 0.5)Ŵ
ŵ
                                 - assume exit loss coefficient = 0.45\Phi- hgate =
\DeltaE.
              if(bigh .gt. dia) bigh = dia
              if (bigh. It. radius) then\mathbf{C}\mathbf CAc = pi/2*(radius**2.-(radius-bigh)**2)beta = a \cos((radians - bigh) / radius)\mathbf{C}\mathbb{C}P = dia * betaendif
C
\mathbf Cif (bigh .gt. radius) then
                 Ac = pi/2.*(radius * *2 + (bigh - radius) * *2)\mathbb{C}^2beta = a cos((bigh-radius)/radius)\mathbf CP = dia * (pi-beta)C
\mathbf Cendif
              if (bigh .eq. radius) then
\mathbf{C}ac = pi/2.*radians*2\mathbf Cp = pi*radius\mathbf Cendif
c endit<br>c if(p.eq. 0.0)then<br>cc write(*,*) 'p = 0.0
                 stop
\mathbf{C}endif
              R = AC/P\mathbf Cif (r \cdot eq. 0.0) then
C
```

```
\mathbf Cwrite (*,*) ' r = 0.0'\mathbf Cstop
           endif
\ddot{\circ}h3 = hdby - elbybC
           if (h3 .1t. 0.0) h3 = 0.0\mathbf Ch12 = k*Cd**2.\mathbf Ch23 = 29. *man**2*L*Cd**2/(R** (4/3))\mathbf Closses = (hdcol - elev) +z-h3-h12-h23-hgate\mathbf Cif (losses .le. 0.0) then\mathbf Cqcol = 0.0C
              else
\mathbf CQcol = Cd*Ac*sqrt(2.*g*losses)\mathbf Cendif
C
ŵ
           endif
C
\mathbf{r}\bulletif (ratio .gt. 1.2 .and. haby .ge. tfour) then\mathbf{C}.
÷
\bulletType 4 culvert flow
\star bigh = hdcol - elev
\starA = pi * radius**2\mathbf CR = dia/4.C
           alpha = bigh+z-hdby-e1byb\mathbf Cif(alpha .le. 0.0)then
C
              qcol = 0.0\mathbf Celse
C
              Qcol = Cd*A*sqrt(2.*g*alpha) / (1.+29.*man**2*L*Cd**2/\mathbf CR***(4/3) +0.45*Cd**2))
\mathbf C\ddot{}endif
\mathbf C\hat{\mathbf{r}}endif
C
\bullet\mathbf{r}if (ratio .gt. 1.2 .and. hdby .lt. tfour) then
C
\Phi\starType 5 culvert flow
\starA = pi*radius**2C
           R = dia/4.\mathbf Calpha = big+z-hdby\mathbb Cif(alpha .le. 0.0)then
C
              qcol = 0.0\mathbf Celse
\mathbf CQcol = Cd*A*sqrt(2.*g*alpha/ha/(1.*29.*man**2*L*Cd**2/C
                     R** (4/3) +0.45*Cd**2))
\mathbf C\divendif
Ć
\starendif
C
\hat{\mathbf{x}}c888 return
C.
         end
\mathbf{A}subroutine opencq hdby hdcolelev slope 1bottomdydxman
\mathbf CIMPLICIT DOUBLE PRECISION (A-H, J-m, O-z)
\mathbf Creal*8 l, man
\mathbf Cparameter (g = 32.2)\mathbf C
```
APPENDIX C

Program Code for PUMPIN.FOR


```
PROGRAM pumpin
         IMPLICIT DOUBLE PRECISION (A-H, J-m, O-z)\mathbf{g}_i\star* polynomial coefficients vol = -1757.08 + 10.7763*hth + 50.5127*hth**2
      derived from the application of the "grapher" best fit 2nd degree
\mathcal{U}\starpolynomial to the Fishman results
\mathbf{w}a=-1757.08b=10.7763c=50.5127\frac{d\mathbf{r}}{d\mathbf{r}}\Phi\begin{array}{cc} \texttt{c50} & \texttt{FORMAT} (\texttt{F5.2}) \\ * & \end{array}\alpha\mathbf{w}♣
         write (*, *) 'Enter the initial head:'<br>write (*, *) ''<br>read (*, *) hi<br>write (*, *) ''
         write (*, *) 'Enter the desired head:'
         write \begin{pmatrix} * & * \end{pmatrix} ''<br>read \begin{pmatrix} * & * \end{pmatrix} '''<br>write \begin{pmatrix} * & * \end{pmatrix} ''
         read(*,*) head
         write (*, *) 'Enter the time period in days:'<br>write (*, *) ''<br>read (*, *) days
         write(*, *) '\Delta\mathbf{w}囊
* convert cfs to acre-ft for time increment dt (days)
* pump = pumping rate in cfs
\frac{d}{d\mathbf{x}}volfinal = a + b*head + c*head**<br>volinit = a + b*hi + c*hi**2volinit = a + b*hi + c*hi**2<br>dvol = volfinal - volinit
         dvol1 = dvol*43560.
\mathbf{x}\starcalculate the pumping rate in cfs for steady state
\frac{1}{2\pi}5O4l67 converts the acre-ft/days to cfs
\pmb{\ast}pump = dvol/days*.5041687\hat{\mathbf{x}}Ŵ
   quadratic equation for smith/bybee lake vol vs elevation in ac ft vs ft mel
\mathcal{A}using the coefficients a, b, c.
\tilde{\mathbf{x}}\mathbf{g}_iwrite(*,10) hi
             write(*, 20) days
             write(*,30) head
             write(*, 40) pump
             write (*, 10) pung<br>write (*, 50) dvol<br>write (*, 60) dvol:
10 format Initial head fS.1 ft
20 format (Time period: 15.1 ft)<br>30 format ('Desired head: 'f5.1 'ft')
```

```
hc = channel head = hdcol - elev\star\bullethc = hdcol - elevC
         if(hc .le. 0.0)then
C.
             qcol = 0.0C
\sigmago to 777
\mathbf Cendif
         if slope .eq. 0.0) then
Ċ.
            write (*, *) 'slope = 0.0'Ċ.
C.
            stop
         endif
\mathbf C\circif (man .eq. 0.0) thenwrite (*, *) 'man = 0.0'C.
C
            stop
d,
         endif
\mathbf{C}if (dy \text{ .eq. } 0.0) then<br>write (*,*)' dy = 0.0'C
            stop
C.
\mathbf Cendif
         if (hdcol .le. elev .or. hdcol .le. hdby) then
C.
            qcol = 0.0C.
\mathfrak{a}goto777
\mathbb{C}endif
\mathbf Cz = slope * LA = (bottom + hc*dx/dy)*hc¢
Ĩ.
         theta = atan(dx/dy)if(cos(theta) .eq. 0.0) then<br>write(*,*) 'cos(theta) = 0.0'\mathbf C\mathbf CÖ.
            stop
         endif
\mathbf CP = bottom + hc/cos(theta)\mathbf CR = A/PC.
         vel = (1.486/man*r**(2/3)*slope**.5\overline{\mathbb{C}}hqate = \text{vel}**2/2/gC
C
         hin = hdcol - hgateif (hin.le. hdby then
C.
             q = 0.0\mathbf CC
         else
             Q = A \star vel\mathbf Cendif
C
\starc777 return
         end
C
\bulletSUBROUTINE INTERT in
           IMPLICIT DOUBLE PRECISION (A-H, J-m, O-z)
           common jns(9000), hns(9000)
                   hby(7500), diffq(7500)C
           real x(9000), y(9000)INTER INTERPOLATES INPUT DATA TO OBTAIN
\mathbf C\mathbf C\mathbf Ctime in Julian day, head in ft MSL, or diffq.<br>variable to interpolate: 1 = hcol\mathbf Ct: time in Julian day, head<br>n: variable to interpolate:
\mathbf C2 = hns\overline{C}3 = diffq\mathbf C4 = hbyd
         returned variable\mathbf Cz:
```
×,

STOP ENI

 \bar{z}

ğ,

APPENDIX

Program Code for HYDRO.FOR

```
this program reads input data and converts it to julian day (1990),
Ŕ
    evaporation and precipitation in inches
      open(unit=10,file='noaamod.dat',status='old')
      open(unit=11, file='hydro.dat', status='unknown')
      read(10, *, end=200) dayj, daym, tair, tdew, prec, wind
10
      tcair = (tair-32)/1.8tcdew=(tdew-32)/1.8rh=( (112.-0.1*tcair+tcdev) / (112+0.9*tcair))**8ea = - 0.132579+0.014123*tair-0.000233125*tair**2+2.98306e-6*tair**3
*convert ea to mm of hg
      ea = ea * 25.4* convert wind speed from mph to mpd
      wind=wind*24
      evap=0.35*ea*(1-rh)*(1+0.0098*wind)*concert mm/d to inch/day
      evap = evap/25.4dayj = dayj + 1096write (11,100) dayj, evap, prec
      format(f5.0, 3x, f8.4, 3x, f4.2)100
      go to 10
200
      stop
```
end

Head (ft MSL)